Impact of Biomass Burning in South and Southeast Asia on Background Aerosol in Southwest China

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ABSTRACT

Biomass burning (BB) in Southeast Asia is particularly pronounced during the dry season. However, the complex topography and long-range transport inherent to Southeast Asia have limited local research on pollution resulting from BB. In this study, the monthly variation in aerosol optical properties at six sites in Southeast Asia (Chiang Mai, Mukdahan, Bac-Lieu, Penang, Singapore, and Bandung) and the fire-point distribution have been analyzed in detail. The Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model was used to simulate the 72-hour back-trajectory from the Shangri-La atmospheric background station in Yunnan Province, China. Our results showed that BB was more common on the Indochinese Peninsula from March to May, whereas it was more common on the Malay Archipelago from August to October due to the latitudinal difference and crop harvest season. Significant BB activity on the Indochinese Peninsula in March resulted in a high surge in extinction (AOD₄₄₀nm = 1.32 ± 0.69, AOD₆₄₇nm = 1.24 ± 0.59) by particles with a smaller diameter (AE = 1.68 ± 0.13) in Chiang Mai. Mapping the long-range transport of BB aerosols reveals that Shangri-La’s pollution was primarily affected by emissions from northern-central India (accounting for 45.2%), and Bangladesh and northern Myanmar (accounting for 38.7%), which indicates that the aerosol pollution on the Yunnan-Guizhou Plateau in springtime could have originated on the southern periphery of the Tibetan Plateau. The results also indicate that BB emission in Southeast Asia had a limited impact on pollution in Southwest China but a relatively large effect on local areas. This study is the first to analyze the trend of aerosols produced from BB in Southeast Asia via ground-based observation, which deepens our understanding of the potential effects of BB aerosols transported long-range from outside Southwest China.

Keywords: Aerosol optical properties; Biomass burning; Southeast Asia; Yunnan-Guizhou Plateau.

INTRODUCTION

Atmospheric aerosols from natural and anthropogenic sources have crucial effects on climate, air quality, and human health (Charlson et al., 1992; Tie et al., 2009; Matus et al., 2012). Li et al. (2017, 2018a) indicated the importance of aerosol-cloud interactions on cloud properties. In addition, climate change can indirectly affect aerosol spatial distribution influenced by total precipitation, as reported by Yang et al. (2010) and Yang et al. (2017). Biomass burning (BB) produces large amounts of aerosol particles, and emissions from BB can be transported widely across a region. Therefore, BB can have a marked impact on global climate (Andreae and Merlet, 2001; Streets et al., 2003; Marlon et al., 2008). BB occurs frequently in the tropics; Andreae et al. (2001) studied the transport of BB aerosols in the equatorial region, finding that they can be transported to the top of the troposphere through deep convection. In Africa, Ansmann et al. (2009) found that dust from the deserts of northern Africa and BB could be transported to South America. In China, Wang et al. (2015a) studied the impact of straw burning on regional aerosol properties in the North China Plain. BB during the dry seasons in Southeast Asia is commonplace, especially around the Indonesian island of Sumatra (Field et al., 2009). The large amount of haze produced during BB can cause severe fine-mode particulate pollution over a short period, and such pollution can harm human respiratory systems (Koe...
et al., 2001). Consequently, BB in Southeast Asia has a marked impact on regional environments, with long-range transportation of aerosol particles to southern China and East Asia in particular attracting scholars’ attention (Chan et al., 2000; Kondo, 2004; Tsay et al., 2016).

The Yunnan-Guizhou Plateau is characterized by a high altitude and low latitude, making it a unique climate (Zheng et al., 2001). Consequently, BB in Southeast Asia has a marked impact on regional environments, with long-range transport making the properties of aerosol in Southwest China differ from those of other areas. The county-level city of Shangri-La is located on the northwestern Yunnan-Guizhou Plateau, connected with the Tibetan Plateau, and lies at an average elevation of 3300 m. The air quality is good under clear conditions, with few sources of local pollution, the majority being from external transport (Zhang et al., 2012a; Gao et al., 2013). Therefore, we can speculate aerosol absorption process may lead to varied aerosol properties in the dry and wet seasons.

Southeast Asia comprises the Indochinese Peninsula and Malay Archipelago. The Indochinese Peninsula (comprising Vietnam, Laos, Cambodia, Thailand, and Myanmar) has a tropical monsoon climate with a high annual temperature. The region has wet and dry seasons, with the wet season running from June to August. Crops are usually planted in the wet season and harvested in the dry season. The Malay Archipelago (including Malaysia, Singapore, and Indonesia) has a tropical monsoon climate with high annual temperatures and adequate rainfall (Manton et al., 2001); rainfall occurs more frequently from November to January. In this region, tropical rainforests are widely distributed and crops are planted year round and can be harvested in all seasons.

Atmospheric aerosol optical depth (AOD) and Ångström exponent (AE) are two basic parameters for studying the optical properties of aerosol (Che et al., 2013; Yu et al., 2015). The atmospheric AOD reflects the degree of extinction of solar radiation by particles across the entire atmosphere, and it is a key parameter for evaluating atmospheric particulate pollution and aerosol radiation effects (Zheng et al., 2018). AE is a vital optical parameter for qualitatively measuring aerosol particle size, and it can be used to initially distinguish aerosol types. Single-scattering albedo (SSA), asymmetry factor, fine-mode aerosol fraction, and particle volume size distributions are also primary parameters of aerosol optical properties (Che et al., 2015b). Numerous studies have demonstrated that BB aerosols make a vital contribution to air pollution in Asia (Deng et al., 2008; Deka and Hoque, 2014). Some studies have determined that aerosol and gases from BB in South and Southeast Asia can be transported to southeastern China and the northwestern Pacific (Jacob, 2003; Zhang et al., 2012b). Aerosol chemical composition data analysis (Wei and Liu, 1994; Streets et al., 2003) has ascertained that BB aerosols can travel far and reach Southwest China. However, due to the large regional span, its complex geographical features, and the lack of ground-based observations in some areas, few studies have directly examined the impact of BB emissions from South and Southeast Asia on air quality and regional climate in Southwest China.

This study is the first to analyze the trend of aerosols produced from BB in Southeast Asia on the basis of ground-based observed AOD and investigate the BB aerosol impact on Southwest China (Che et al., 2014). In order to consider topographical factors (Venkataraman et al., 2005; Engling et al., 2011), we choose the Shangri-La (Yunnan Province, China) atmospheric background station, which can be used to represent the changes of background aerosols in Southwest China, combined with European Centre for Medium-Range Weather Forecasts (ECMWF) wind field data, the Moderate Resolution Imaging Spectroradiometer (MODIS) satellite data, and particulate pollution 72-hour back-trajectory modeling. The findings are scientific significance of the sources of spring pollutants in Southwest China and changes in aerosol optical properties during the biomass burning in Southeast Asia. The remainder of this paper is structured as follows: “Introduction” summarizes the research on BB, the climatic conditions of the Yunnan-Guizhou Plateau and Southeast Asia, and the challenges of investigating BB in Southeast Asia. “Data and methodology” introduces the research location, observation instruments, and data source. In Section 3, the monthly aerosol optical properties in Southeast Asia are analyzed. In Section 4, the aerosol optical properties and sources of spring pollutants in Shangri-La are analyzed. Finally, Section 5 presents the discussion and conclusions.

**DATA AND METHODOLOGY**

**Site Information**

The county-level city of Shangri-La is located in the Diqing Tibetan Autonomous Prefecture of Yunnan Province, China, and it borders the provinces of Sichuan and Tibet. Shangri-La is located on the eastern edge of the Tibetan Plateau; it is east of the Sanjiang area’s longitudinal valley and to the southeast of the Hengduan Mountains. The overall trend of the geography is northwest high and southeast low, with an average altitude of 3459 m (Fig. 1). The atmospheric background observation station (28.01°N, 99.44°E) used for data collection in this study is located on the top of a hill 12 km northeast of Shangri-La. The observation field is 3583 m above sea level and was completed in 2009 (Wang et al., 2015b). It is one of the key atmospheric background observation stations in China and a member of the World Meteorological Organization’s Global Atmosphere Network. The atmospheric background observation station in Shangri-La is the most advanced in China; thus, it can accurately measure the background atmospheric environment in Southwest China and reflect
As shown in Fig. 1, based on the latitude distribution (from north to south), long-term data observation and dense fire-point areas, the six selected observation sites in Southeast Asia were as follows: Chiang Mai, Thailand (18.77°N, 98.97°E); Mukdahan, Thailand (16.61°N, 104.68°E); Bac-Lieu, Vietnam (9.28°N, 105.73°E); Penang, Malaysia (5.36°N, 100.30°E); Singapore (1.30°N, 103.78°E); and Bandung, Indonesia (6.89°S, 107.61°E). Detailed information regarding these locations is presented in Table 1.

Data and Methodology

The aerosol optical properties in this study were derived from AERONET (http://aeronet.gsfc.nasa.gov/), which is developed and supported by the U.S. National Aeronautics and Space Administration, using the direct radiation observation data and scattered radiation observation data (Holben et al., 1998; Dubovik et al., 2000). The AOD used in this study was version 3 of the Direct Sun Algorithm with Level 1.5 Cloud Filtering. Shangri-La data comes from the China Aerosol Remote Sensing Network (CARSNET) using CE318 sun photometers (Che et al., 2009). This is an automatic multispectral sun tracking photometer, designed for accurate sun measurements (Che et al., 2008). Sun photometers are basic instruments used for observing the optical properties of the atmosphere (Estellés et al., 2012). Employing eight observation channels measured direct solar radiation data using the CE318 sun photometer can be used to inversely calculate atmospheric transmittance, AOD, total amount of atmospheric water vapor, and total amount of ozone. Moreover, CE318 sun photometers can be used to survey the entire columnar atmosphere in the vertical direction, obtain column-integrated aerosol optical properties, and verify aerosol parameters retrieved by satellites (Dubovik and King, 2000).

We applied the Hybrid Single Particle Lagrangian Integrated Trajectory (HYPLIT) model (Stein et al., 2015) in MeteoInfo (Wang, 2014) to conduct 72-hour back-trajectory analysis from March 1, 2012, to May 30, 2012, in Shangri-La. The clustering mode was employed to cluster the main back trajectory according to the velocity and direction of the air mass.

Wind field data were derived from reanalysis of the ECMWF (http://apps.ecmwf.int/datasets/) images, which display areas that were burning at the time of overpass. The MODIS electromagnetic spectrum distribution is 0.4–0.14 nm, with resolutions of 250, 500, and 1000 m, and acquires global data four times per day; therefore, it can accurately monitor forest fire points (Giglio, 2013).

RESULTS AND DISCUSSION

Monthly Variation in Aerosol Optical Properties in Southeast Asia

Variation in AOD

In this section, the changes in daily average AOD data at the six selected stations in Southeast Asia (Chiang Mai, Mukdahan, Bac-Lieu, Penang, Singapore, and Bandung) are analyzed. As shown in Figs. 2(a)–2(b), AOD at 440 nm in the six areas was larger than AOD at other wavelengths, and the optical depth decreased with increasing wavelength in accordance with the variation in AOD in urban aerosol (Eck et al., 1999). In mainland Asia, anthropogenic aerosol particle emissions tend to exhibit greater wavelength-selective attenuation of sunlight than do larger particles.
Table 1. The information of stations in Southeast Asia.

<table>
<thead>
<tr>
<th>Station</th>
<th>Country</th>
<th>Latitude</th>
<th>Longitude</th>
<th>Altitude (m)</th>
<th>Starting Time</th>
</tr>
</thead>
<tbody>
<tr>
<td>Chiang Mai</td>
<td>Thailand</td>
<td>18.77°N</td>
<td>98.97°E</td>
<td>312.0</td>
<td>2006</td>
</tr>
<tr>
<td>Mukdahan</td>
<td>Thailand</td>
<td>16.61°N</td>
<td>104.68°E</td>
<td>166.0</td>
<td>2003</td>
</tr>
<tr>
<td>Bac-Lieu</td>
<td>Vietnam</td>
<td>9.28°N</td>
<td>105.73°E</td>
<td>10.0</td>
<td>2003</td>
</tr>
<tr>
<td>Penang</td>
<td>Malaysia</td>
<td>5.36°N</td>
<td>100.30°E</td>
<td>51.0</td>
<td>2011</td>
</tr>
<tr>
<td>Singapore</td>
<td>Singapore</td>
<td>1.30°N</td>
<td>103.78°E</td>
<td>30.0</td>
<td>2006</td>
</tr>
<tr>
<td>Bandung</td>
<td>Indonesia</td>
<td>6.89°S</td>
<td>107.61°E</td>
<td>826.0</td>
<td>2009</td>
</tr>
</tbody>
</table>

Fig. 2. Box plots depicting monthly variation of AOD at 440 nm, 675 nm, 870 nm, and 1020 nm wavelengths at six stations in Southeast Asia: (a) Chiang Mai, Mukdahan, and Bac-lieu and (b) Penang, Singapore, and Bandung. The central line in each box is the median, and the lower and upper limits are the first and third quartiles, respectively. The lines extending vertically from the box indicate the spread of the distribution with the length being 1.5 times the difference between the first and third quartiles. The square symbols indicate the geometric means.
such as dust aerosol (Che et al., 2015a). Therefore, this study focused on AOD changes at 440 nm. In sunny conditions, $\text{AOD}_{440nm}$ should be less than 0.50; $\text{AOD}_{440nm}$ greater than 1.0 indicates poor air quality and polluted weather with abundant particulates.

Chiang Mai, Mukdahan, and Bac-Lieu are all located on the Indochinese Peninsula and exhibited approximately equal AOD changes (Fig. 2(a)): From January to March, $\text{AOD}_{440nm}$ increased significantly, peaking in March (Chiang Mai: $1.32 \pm 0.69$; Mukdahan: $0.98 \pm 0.38$); this finding is consistent with the peak in $\text{AOD}_{500nm}$ ($\text{PM}_{2.5} = 120 \mu g \cdot m^{-3}$) on April 3, 2014, in Chiang Mai (Pani et al., 2018). From March to May, $\text{AOD}_{440nm}$ decreased gradually but remained above the mean value. $\text{AOD}_{440nm}$ Values were lowest during the wet season (June–August) before markedly increasing from September to November. This trend was most obvious in Thailand, with two peaks in the monsoon rainfall season. This finding is consistent with the results of Gautam et al. (2013). Overall, $\text{AOD}_{440nm}$ on the Indochinese Peninsula varies depending on the timing of the rainy season and thus follows a strong seasonal cycle. However, the overall value at Bac-Lieu was lower than at the other two stations, indicating higher air quality.

Penang, Singapore, and Bandung are located on the Malay Archipelago and all exhibited approximately equivalent changes (Fig. 2(b)): $\text{AOD}_{440nm}$ Varied little from...
January to July and increased significantly from August to October; however, only the October AOD\textsubscript{440nm} average value of Penang (1.07 ± 1.36) was higher than 1.0. AOD\textsubscript{440nm} values were lowest during the wet season (November–January). Notably, the three areas’ variation tendencies differed: Changes in AOD\textsubscript{440nm} in Penang were relatively flat in August and September but more prominent in October, AOD\textsubscript{440nm} in Singapore was highest in September and October, and AOD\textsubscript{440nm} fluctuation in Bandung was low from January to October.

Variation in Fine-mode AOD

The total aerosol optical depth (AOD-T) can be further divided into fine-mode particle extinction optical depth (AOD-F) and coarse-mode particle extinction optical depth (AOD-C). The distribution of aerosols from a 1997 Indonesian forest fire indicated that BB produces numerous fine particles, known as black carbon aerosol (Narukawa et al., 1999). Whether the fine particulate matter is the dominant particle in aerosol can be determined by analyzing the proportion of fine-mode particles in the total AOD.

Fig. 3 illustrates that in Chiang Mai, Thailand, the maximum FAOD\textsubscript{440nm} was 1.24 ± 0.59 in March, and the minimum was 0.12 ± 0.06 in August. After the rainy season, FAOD\textsubscript{440nm} fluctuated from September to November but peaked in October (0.44 ± 0.23), indicating that fine-mode particles dominated from February to May with another fine-mode particle increase from September to November. Changes in the Mukdahan area largely mirrored those in Chiang Mai. In Bac-Lieu, Penang, and Singapore, the fine-mode particles exhibited a significant increase from August to October, with peaks occurring in mostly September or October (Bac-Lieu: 0.30 ± 0.29 in September; Penang: 0.758 ± 0.80 in September, 0.756 ± 0.89 in October; Singapore 0.87 ± 0.57 in October). However, the overall FAOD\textsubscript{440nm} value (0.25 ± 0.07) in Bac-Lieu was low; combined with AOD\textsubscript{440nm} (0.30 ± 0.08), fine-mode particles still dominated in this area’s air pollution.

Variation in AE

AE is often employed as a qualitative indicator of aerosol particle size. AE interpolated by spectral AOD at 440 and 870 nm (AE\textsubscript{440-870nm}) is more accurate, and thus AE\textsubscript{440-870nm} is used to indicate change trends in the Ångström wavelength index. We mainly focus on AE\textsubscript{440-870nm} > 1.0 because larger AE values represent an overall smaller atmospheric particle size, and values less than 1.0 indicate large particles such as sea salt and dust. If the average AE\textsubscript{440-870nm} value is greater than 1.0, it indicates that most particles in the atmosphere are small particles associated with combustion (Schuster et al., 2006). As shown in Fig. 4, the monthly average values of AE\textsubscript{440-870nm} in all six regions were greater than 1.0, indicating that fine particles in the atmosphere dominated as pollutants compared with large particles such as dust and sea salt.

The stations on the Indochinese Peninsula exhibited a similar pattern, with an upward trend from January to March, followed by a rapid decrease from March to June, a slight increase from June to September, and little change from September to December. However, the trends vary slightly by location. The monthly average value of AE\textsubscript{440-870nm} in Chiang Mai was largest (1.68 ± 0.13) in March and lowest in June (0.92 ± 0.34), with a subsequent upward trend. By contrast, the average value of AE\textsubscript{440-870nm} in Mukdahan was highest in March (1.59 ± 0.12), and the average value of AE\textsubscript{440-870nm} had an upward trend after June, with a small peak in September (1.31 ± 0.31) and no significant change from October to December. Bac-Lieu demonstrated the

Fig. 3. Box plots of monthly fine-mode particle AOD at a wavelength of 440 nm across the six areas (Chiang Mai, Mukdahan, Bac-Lieu, Penang, Singapore, and Bandung). In each box, the central blue line is the median, and the lower and upper limits are the first and third quartiles, respectively. The red square symbols indicate the geometric means.
Fig. 4. Box plots of monthly AE properties at a wavelength of 440 nm in the six areas (Chiang Mai, Mukdahan, Bac-Lieu, Penang, Singapore, and Bandung).

largest value of $\text{AE}_{440\text{-}870\text{nm}}$ in April and the minimum value in July, with another slight increase from July to September that gradually decreased from September to December. The trend of $\text{AE}_{440\text{-}870\text{nm}}$ in Singapore was in agreement with that found by Salinas et al. (2009), with April to October dominated by mostly fine-mode particles with few coarse-mode particles.

**Variation in SSA**

Atmospheric aerosol SSA is used to characterize the absorption and scattering of atmospheric aerosols, which is defined as the percentage of total extinction that is aerosol scattering extinction. SSA size is determined mainly by the shape, composition, and concentration of atmospheric aerosol particles, and it is a crucial parameter that directly reflects the scattering intensity of atmospheric aerosol particles. Moreover, it is a key variable for evaluating the effect of aerosol on climate. According to Dubovik et al. (2002), the average value of $\text{SSA}_{440\text{nm}}$ in the Northern Hemisphere is 0.85–0.95. A monthly mean $\text{SSA}_{440\text{nm}}$ higher than 0.85 indicates that the aerosol particles in an area have a high degree of scattering. By contrast, if it is less than 0.85, the aerosol particles in an area are more absorptive.

Fig. 5 reveals that the monthly average of $\text{SSA}_{440\text{nm}}$ in the six locations in Southeast Asia fluctuated between 0.85 and 0.95, indicating that the aerosol particles were more scattered. However, the fluctuation at the Chiang Mai station is clearer: The largest $\text{SSA}_{440\text{nm}}$ was in May (0.91 ± 0.30), and the minimum was in August (0.82 ± 0.07). $\text{SSA}_{440\text{nm}}$ began increasing after the wet season, and the second highest value appeared in October (0.90 ± 0.06). In Chiang Mai, the monthly mean of $\text{SSAT}_{440\text{nm}}$ was 0.87 ± 0.03, which was markedly lower than those of the other five regions. Li et al. (2013) analyzed the SSA of aerosols in Phimai, Thailand, finding that they ranged from 0.86 ± 0.04 to 0.92 ± 0.02, which are similar values to those recorded in Chiang Mai in the present study. These findings demonstrate that aerosol particles in Thailand are absorptive and have a low degree of scattering.

**Aerosol Particle Volume Size Distributions**

Atmospheric aerosols comprise many particles of different sizes, also called multispectral aerosols. Volume size distributions are key physical indexes that describe the distribution of particle size concentrations for multispectral aerosols. Particle size concentrations are among the decisive factors in determining the transport characteristics, life expectancies, and optical properties of atmospheric aerosols. For irregularly shaped particles, volume size most effectively expresses the effect of particles on extinction, so the volume is taken as the characteristic dimension parameter in the particle spectral distribution function.

Fig. 6 shows that bimodal distributions were present in all six regions. The aerosol fine particles were mainly 0.1–0.3 μm, and the coarse particles were mainly 5–10 μm. In March and April, the concentration of fine aerosol particles in Chiang Mai and Mukdahan increased significantly, which may be related to local BB, whereas the coarse-mode particle concentrations were greater in the rainy season (July–September), which may be related to gas-particle conversion under high temperature and humidity conditions. In Mukdahan, the percentage of particles with radii from 0.5 to 5 μm was significantly higher in July than it was in other months, which may be related to the increase in some aerosol particles’ hygroscopicity during the rainy season.

On the Indochinese Peninsula, concentrations of fine aerosol particles were significantly higher from July to September, which may be related to the incineration of tropical plants (especially Musaceae) in Indonesia, which produces substantial fine-grained haze and takes time to...
reach surrounding areas. Coarse-mode particle concentrations were higher in the rainy season (November–January), and the radii of the particles were mostly 5 µm; the concentrations of coarse-mode particles were particularly high in November and December.

**Analyzing Absorptive Atmospheric Aerosol Changes**

In this section, the statistical inversion data regarding absorption AOD (AAOD) at 440 nm (AAOD$_{440\text{nm}}$) and absorption AE at 440–870 nm (AAE$_{440-870\text{nm}}$) are examined. Absorptive aerosols lead to the redistribution of solar radiation in the vertical direction and directly affect the elimination of the physical processes of atmospheric convective clouds, indirectly affecting the global and regional climates. Absorptive aerosols reduce solar radiation and prevent it from reaching the ground, increasing atmospheric absorption and backscatter. Concurrent absorption of surface anti-radiation by absorptive aerosols compensates for backscattered radiation in the top of the atmosphere. Therefore, the radiative forcing of absorptive aerosols is generally less than that of purely scattering aerosols (Zhu et al., 2014).

Atmospheric aerosol extinction consists of absorption and scattering. Absorption AE (AAE) is calculated from spectral absorption AOD (AAOD) at 440 and 870 nm, and AAOD is calculated from AOD and SSA by Eqs. (1) and (2) (Dubovik et al., 2002):

\[
\text{AAOD} (\lambda) = \text{AOD} (\lambda) \times [1 - \text{SSA} (\lambda)] 
\]

(1)

\[
\text{AAE} = -\frac{\text{dln}[\text{AAOD} (\lambda)]}{\text{dln}(\lambda)}
\]

(2)

BB produces substantial black carbon, and black carbon is a mainly absorptive aerosol (Jacobson, 2001). Fig. 7 reveals that AAOD$_{440\text{nm}}$ varies significantly from January to May in Thailand, especially in Chiang Mai (January: 0.051 ± 0.02; February: 0.10 ± 0.05; March: 0.17 ± 0.08; April: 0.12 ± 0.06; May: 0.053 ± 0.02), whereas AAOD$_{440\text{nm}}$ in Mukdahan was higher in February and March. After the wet season, a small increase in AAOD$_{440\text{nm}}$ occurred in Thailand. Overall AAOD$_{440\text{nm}}$ varied little in other areas; in Vietnam and Malaysia, a slight increase occurred from August to October, and in Singapore and Indonesia, a slight growth occurred from September to November.

According to Eck et al. (2010), the AAE value of aerosol particles mainly composed of absorptive black carbon aerosol is approximately 1.0. As organic particles in the aerosol increase, the AAE value increases between 1.0 and 2.0; by contrast, the AAE of BB aerosol is between 1.2 and 1.5.

Changes in AAE$_{440-870\text{nm}}$ are shown in Fig. 8. Similar to the trend of AAOD$_{440\text{nm}}$ areas in Thailand exhibited marked variation, fluctuating from January to May with a maximum AAE$_{440-870\text{nm}}$ in March (Chiang Mai: 1.35 ± 0.21; Mukdahan: 1.24 ± 0.22). The distribution of AAE$_{440-870\text{nm}}$ in Thailand is large, which indicates that background absorptive aerosols in the region comprise mainly BB aerosol particles and some sand dust-type aerosol particles. No obvious change in the annual average of AAE$_{440-870\text{nm}}$ was observed in other areas (Bac-Lieu: 0.79 ± 0.08; Penang: 0.97 ± 0.11; Singapore: 0.87 ± 0.13; and Bandung: 0.85 ± 0.12). AAE$_{440-870\text{nm}}$ in Penang was significantly higher than it was in other regions; however, the annual average values across all regions were less than 1, probably because black carbon aerosol in the region was coated with nonabsorptive material (Jacobson, 2000).

**Analysis of Aerosol Optical Properties and Potential Sources in Shangri-La**

According to the analysis of the aerosol optical properties at the six stations in Southeast Asia presented in
Section 3, BB produced a large amount of fine-mode aerosol in Thailand from March to May, which is evident from the aerosol optical properties. Although burning of tropical plants is widespread in Indonesia in July and August, because of the large area, long transmission range, general atmospheric circulation, topography, and other factors, this burning is less obviously reflected in the aerosol optical properties. In Section 4, the main sources responsible for aerosol optical properties and pollution in Shangri-La from March to May are examined. In addition, the influence of BB emission in South Asia and Southeast Asia on background aerosol in Southwest China is discussed.

**Monthly Variation in Aerosol Optical Properties in Shangri-La**

As shown in Fig. 9 and Table 2, according to the AOD distribution, AOD$_{440\text{nm}}$ has the largest overall value, and AOD decreases as the wavelength increases, which conforms to the trend of urban-type aerosol in mainland Asia (Qi et al., 2016). The annual average AOD$_{440\text{nm}}$ in this area is $0.11 \pm 0.05$, indicating excellent air quality. AOD$_{440\text{nm}}$ values in March–May and July–September are significantly higher than the annual mean, with the maximum appearing in August ($0.19 \pm 0.02$), indicating that AOD$_{440\text{nm}}$ in Shangri-La is higher in spring and summer and lower in autumn and winter. AOD$_{440\text{nm}}$ was particularly high in the spring (March: $0.17 \pm 0.09$; April: $0.13 \pm 0.07$; and May: $0.11 \pm 0.07$), with values all above the mean value, although a downward trend was observed.

The annual average AE$_{440-870\text{nm}}$ in Shangri-La was $1.26 \pm 0.14$, which reveals that the atmospheric aerosols contained more fine particles with fewer coarse particles such as sand dust from northwestern China. The maximum average AE$_{440-870\text{nm}}$ appeared in October ($1.50 \pm 0.44$), and the minimum appeared in August ($1.08 \pm 0.29$). No obvious fluctuation was exhibited from March to May, but a sharp
fluctuation was noted from June to October; these findings indicated that most pollutants were fine particles that are more prevalent during the spring. The values for $\text{AE}_{440-870\text{nm}}$ in March (1.20 ± 0.37), April (1.20 ± 0.33), and May (1.09 ± 0.41) demonstrated that the trend of $\text{AE}_{440-870\text{nm}}$ for this period was the same as that of $\text{AOD}$, with fine particles in the atmosphere dominating. However, $\text{AE}_{440-870\text{nm}}$ was slightly lower in May, which may be because of more rainfall in May and increased humidity leading to hygroscopic growth in aerosol. Alternatively, more dust may have been transported from the Middle East; either of these reasons could cause a decline in the mean value of $\text{AE}_{440-870\text{nm}}$.

**Analysis of the Sources of Pollutants in Shangri-La from March to May**

Particle back-trajectory analysis of the Shangri-La region in March to May 2012 (Fig. 10) revealed that pollutants over the 72-hour period were primarily from northern Indian, Bangladesh, and northern Myanmar and were transmitted.
Fig. 9. Box plots of monthly variation in aerosol optical depth at wavelengths of 440, 675, 870, and 1020 nm in Shangri-La. Double-Y ordinates: Left axis shows AOD and right axis shows AE (440–870 nm). The black square symbols indicate the monthly mean of AE440-870nm.

Table 2. Comparison of aerosol optical parameters including AOD at 440 nm, 675 nm, 870 nm, 1020 nm, and AE440-870nm in Shangri-La.

<table>
<thead>
<tr>
<th>Month</th>
<th>AOD440nm</th>
<th>AOD675nm</th>
<th>AOD870nm</th>
<th>AOD1020nm</th>
<th>AE440-870nm</th>
</tr>
</thead>
<tbody>
<tr>
<td>Jan</td>
<td>0.06</td>
<td>0.04</td>
<td>0.03</td>
<td>0.03</td>
<td>1.16</td>
</tr>
<tr>
<td>Feb</td>
<td>0.08</td>
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Fig. 10. Topographic map and 72-hour back trajectories with three types of cluster at the Shangri-La site during March 2012.
along the southern periphery of the Tibetan Plateau. Indo-Asian pollution has a great influence on Southwest China (45.16%), which is consistent with previous research on the great Indo-Asian haze (Ramanathan et al., 2001). Due to the uplift area of the Tibetan Plateau, potential pollution sources from as far as the Middle East and North Africa can be transported to Southwest China (Xin et al., 2016). However, apart from northern Myanmar, few trajectories are from Southeast Asia, and therefore pollution from BB in northern Thailand has limited effect on Shangri-La.

The monthly average wind direction analysis is presented in Fig. 11. The wind direction varied little from
March to May, which was dominated by a southwesterly wind at the ground level, with a horizontal westerly wind at higher levels. Wind speed was low but increased with altitude. The wind field at 850 hPa was dominated by a southwesterly wind, indicating that pollution in northern Myanmar could easily be delivered to Shangri-La but diminishing the possibility of pollution transport to Shangri-La from northern Thailand. Westerly winds prevailed at 700 and 500 hPa, indicating that pollution from Bangladesh, northern India, and even the Middle East could be transmitted to Southwest China. At the same time, pollutants in Southeast Asia can also be transported downwind efficiently to the Guangxi and Guangdong regions in China.

Fig. 12 presents satellite images of potential fire points; the red points on the photos are fire points, and the gray zones are haze. Large areas of Southeast Asia are gray with many unnumbered fire points across the Indochinese Peninsula, indicating that BB is widespread throughout the region and responsible for haze that causes heavy air pollution. In early March, BB is concentrated mainly in the central and southern Indochinese Peninsula, such as in Thailand, Laos, and Vietnam. From there, BB gradually moves northwestward. By the end of March, BB is primarily in Myanmar, Bangladesh, and northern India.

CONCLUSIONS

In this study, using ground observation data from Southeast Asia (Chiang Mai, Mukdahan, Bac-Lieu, Penang, Singapore, and Bandung) to assess BB emissions and their impact on local atmospheric aerosol optical properties. By combining wind-field, fire-point, and back-trajectory analysis (using the HYSPLIT model), the aerosol optical properties and potential sources of pollution in Shangri-La
were analyzed and discussed with a focus on the impact of BB pollution from northern Thailand and Myanmar in spring. The potential effects of long-range-transported BB aerosols on the study area were evaluated, and the main sources of emissions arriving in Southwest China were identified. The major conclusions from the results are summarized as follows:

(1) Due to the large latitudinal range and different climatic factors in Southeast Asia, the prevalence of BB differs by nation. Based on the analysis of aerosol optical properties in Southeast Asia, countries located on the Indochinese Peninsula tend to conduct large-scale BB activities from March to May of each year. Following the wet season (September–November), small-scale BB may occur in some areas, or BB in Indonesia and other regions may emit large amounts of fine particulate matter that are transmitted to the Indochinese Peninsula. By contrast, countries and regions located on the Malay Archipelago have no obvious and frequent incidents of BB. However, the period of August to November (prior to the wet season) is deeply affected by BB emissions from Indonesia’s southern island. Moreover, some small-scale BB activities may occur in other regions.

(2) Whereas BB emissions from Thailand have only a slight impact on pollution in Shangri-La, those from northern Myanmar during the spring clearly affect this city. Furthermore, although BB emissions from other parts of Southeast Asia have a limited impact on Shangri-La, pollutants can be transported downwind efficiently to China.

(3) The pollution in Shangri-La from March to May is primarily contributed by sources along the pollution conveyor belt on the southern periphery of the Tibetan Plateau, including anthropogenic emissions from northern-central India, Bangladesh, and northern Myanmar, which account for 45.16% of the total external pollution. Pollution from the Middle East and northern Africa can also be transported to Shangri-La. Therefore, springtime background aerosol pollution on the Yunnan-Guizhou Plateau comes mainly from South Asia and northern Myanmar (in Southeast Asia). Although BB in Southeast Asia is not the main external source of pollutants in Southwest China, it has a notable local impact on the Indochinese Peninsula and Malay Archipelago.

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