

66 dependence of AAE can be used to provide a quantitative estimate of BrC (Sandradewi et al.,
67 2008; Harrison et al., 2012), and also for further distinction of sources (Ajtai et al., 2011; Pinter
68 et al., 2018). However, theoretical calculations and field studies have shown that the likely
69 range of AAE for internally mixed BC varies in a range of approximate 1-1.7, depending on
70 the size and optical properties of the particle core and non-absorbing coating and the
71 wavelength pairs used to determine AAE (Gyawali et al., 2009; Lack and Cappa, 2010; Lin et
72 al., 2016). Using a three wavelength photo-acoustic soot spectrometer (PASS-3) and aerosol
73 mass spectrometers (AMS), Yuan et al. (2016) explored the relationship between the measured
74 AAE and the relative abundance of organic aerosol to BC and derived more realistic AAE
75 values for “pure” BC aerosol, which is significantly lower than 1. Contributions to total light
76 absorption at 370 nm by BrC were also estimated by comparing the wavelength dependence
77 of the absorption coefficients of aerosols before and after heated (Guo et al., 2014). This is
78 based on the assumption that the change of AAE for aerosols after heated is due to the
79 evaporation of BrC, but the potential error of which has not been fully assessed.

80 In the present work, the optical, physical, and chemical properties of atmospheric aerosols
81 were measured before and after heated in the Pearl River Delta (PRD) region, China. In this
82 region, organics are the second largest contributor for light extinction (Zhang et al., 2013a;
83 Chen et al., 2016), and the average light absorption contributions of BrC at 405 nm were
84 quantified to be ~10%. This work presents spectral measurements of light absorption
85 coefficient (σ_{abs}) in a broad spectral range (370–950 nm) by an Aethalometer, highlighting
86 some characteristics of light-absorbing aerosols. An improved method was developed, based

87 on wavelength segregated AAE, in order to estimate of the contribution of BrC to light
88 absorption at 370 nm ($\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$). Meanwhile, the BrC contribution was also compared
89 to those estimated based on the previously reported methods.

90

91 **2 Methods**

92 **2.1 Ambient observation**

93 Ambient measurements were performed at Guangdong Atmospheric Supersite (22.73N,
94 112.93E), a suburban site (Zhang et al., 2014) in the PRD region from July 8 to August 2,
95 2014. Over the sampling period, ambient temperature and relative humidity generally varied
96 between 24–37 °C and 40–93%, with an average of 29 °C and 73%, respectively. The
97 concentration of PM_{2.5} varied in a wide range of 0.78–166 $\mu\text{g m}^{-3}$, with mean value of 29.7 μg
98 m^{-3} . Wind speed ranged between 0.3–5.1 m s^{-1} and the prevailing direction was from
99 southwestern. Wind speed was lower ($\sim 1 \text{ m s}^{-1}$) during the early morning hours and was
100 relatively higher ($\sim 2 \text{ m s}^{-1}$) in the afternoon.

101 To compare the properties of aerosols before and after heated, an automatic-controlled
102 actuated ball valve that switches between “bypass” and “heating” on a 5-min cycle. The
103 thermodenuder (TD, Dekati Ltd.) consists of a 95-cm long stainless steel tube, where the
104 particles can be heated with controlled temperature, i.e., 250 °C. The residence time of
105 particles in the TD section is estimated to be 1.2 s, which should be enough for the effective
106 vaporization of semi-volatile species (Villani et al., 2007). Volatile or semi-volatile species,
107 such as most of organic compounds and ammonium nitrate, are supposed to evaporate during

108 the heating (Bi et al., 2015). Subsequently, these species were adsorbed in active charcoal in
109 the adsorber section. The bypass consists of a silica tube of similar length for heating condition.
110 In order to eliminate the influence of resident aerosols in the tubes, only the measurements
111 during the latter 3-min of the 5-min cycle were included in the following data analysis. Particle
112 loss could be typically represented as a function of flow rate and particle size (Stevanovic et al.,
113 2015). The test made by the manufacturer shows that particles with a physical diameter $d_p \leq$
114 70 nm have a size dependent particle loss, increasing from ~20% to ~40% when d_p decreases
115 from 70 nm to 10 nm; particle loss at $d_p > 70$ nm is estimated to be approximately 20% at the
116 total flow rate of 10 lpm, mainly attributed to the diffusion of particles. Regarding that the d_p of
117 refractory BC was reported to be around 200 nm in the PRD region (Huang et al., 2011), the
118 particle loss might lead to a 20% negative bias for σ_{abs} measured through the heated line. The
119 BC around 200 nm was most probably attributed to be from traffic emission, with their diurnal
120 variation similar to that of NO_x (Qin et al., 2017), as shown in Fig. S1.

121 The σ_{abs} was determined in 1-min resolution by an Aethalometer (AE-33, Magee
122 Scientific) with seven wavelengths (at 370, 470, 520, 590, 660, 880 and 950 nm, respectively).
123 A single particle aerosol mass spectrometer (SPAMS, Hexin Analytical Instrument Co., Ltd.)
124 was deployed to measure the size and mass spectra of individual particles (Li et al., 2011).
125 Aethalometer and SPAMS shared an aerosol inlet (dried total suspended aerosols) and were
126 placed downstream of the TD. Additionally, the concentrations of water soluble ions
127 associated with $\text{PM}_{2.5}$ were obtained from a Monitor for AeRosols and GAses in ambient air
128 (MARGA) (Rumsey et al., 2014). Hourly variations of organic carbon (OC)/elemental carbon

129 (EC) in PM_{2.5} were provided by a Sunset Semi-Continuous Aerosol Analyzer. Concentration of
130 O₃ was measured using an ozone analyzer (Model 49i, Thermo-Fisher Scientific). For
131 comparison, all the measured parameters were present as hourly averages in this paper. A
132 detailed description of the instruments and the data analysis can be found in Supporting
133 Information (SI).

134

135 **2.2 Light absorption measurements**

136 Generally, filter-based optical measurement instruments (e.g., Aethalometer) must take
137 into account the “filter loading effect” to accurately determine the σ_{abs} (Arnott et al., 2005;
138 Ajtai et al., 2011). The new version Aethalometer (AE-33) used in the present study has been
139 improved with the filter loading correction part, based on a two parallel spot measurement of
140 optical absorption, through providing a real-time output of the ‘loading compensation’
141 parameter to compensate for the “loading effect”, in detail described in the SI. The details of
142 operation principle, data deduction, error budget of AE-33, inherent uncertainties in its
143 technique and the corrections are extensively available in the literature (Drinovec et al., 2015).
144 The BC mass concentration is calculated from the change in optical attenuation at 880 nm in
145 the selected time interval using the mass absorption cross section (MAC) of 7.77 m² g⁻¹,
146 which is close to that reported in the PRD region (Cheng et al., 2011). The noise level of
147 AE-33 at the time-base of 1 minute is < 0.016 Mm⁻¹ for the σ_{abs}
148 (<http://www.magescientific.com/>). The contribution of dust to σ_{abs} can be considered to be

149 negligible in the PRD region (Yuan et al., 2016), due to limited mass fraction (Cheng et al.,
150 2008; Zhang et al., 2013a), and lower mass absorption efficiency values (Favez et al., 2009).

151

152 **3 Results and Discussion**

153 **3.1 Previous methods to estimate the light absorption from BrC**

154 **3.1.1 Method I: assuming AAE = 1 for “pure” BC**

155 The wavelength (λ) dependence of σ_{abs} is usually approximated by a power-law expression
156 (e.g., Bergstrom et al., 2007), defined as:

$$157 \quad \sigma_{\text{abs}} = K \times \lambda^{-\text{AAE}} \quad (\text{R1})$$

158 where K is a constant. In practice for multi-wavelength measurements, AAE can be
159 retrieved by fitting the σ_{abs} to the λ (370–950 nm), or to the different wavelength ranges
160 (Bergstrom et al., 2007). Thus, a wavelength segregated AAE value can be obtained through
161 the logarithmic fit of R1, or $\text{AAE} = \ln(\sigma_{\text{abs}1}/\sigma_{\text{abs}2})/\ln(\lambda_1/\lambda_2)$ for two wavelengths.

162 The AAE has often been used as a simplified method for attributing short visible
163 wavelength absorption to BrC (Lack and Cappa, 2010). Light absorption at longer wavelength
164 (950 nm) is assumed to be due only to BC, and is extrapolated to the shorter wavelength using
165 an assumed AAE, most commonly to be one, which is illustrated in Fig. 1. It estimates the
166 $\sigma_{\text{abs,BrC}}$ at 370 nm by subtracting the measured σ_{abs} to the extrapolated $\sigma_{\text{abs,BC}}$ at 370 nm
167 (“Method I”):

$$168 \quad \sigma_{\text{abs,BrC}} = \sigma_{\text{abs,370nm}} - \sigma_{\text{abs,950nm}} \times (370/950)^{-\text{AAE}} \quad (\text{R2})$$

169

170 **3.1.2 Method II: based on light absorption enhancement**

171 The $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ was also estimated based on the enhancements of σ_{abs} (E_{abs}), defined
172 as the ratio of σ_{abs} for ambient and heated aerosols ($E_{\text{abs}} = \sigma_{\text{abs,ambient}}/\sigma_{\text{abs,TD}}$), at both 370 and 950
173 nm in previous studies (Cappa et al., 2012; Guo et al., 2014).

$$174 \quad \sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}} = (E_{\text{abs,370nm}} - E_{\text{abs,950nm}})/E_{\text{abs,370nm}} \quad (\text{R3})$$

175 where $\sigma_{\text{abs,BrC}}$ refers to the effective absorption coefficient at 370 nm for the BrC
176 evaporating under the TD temperature (i.e., 250 °C), which is slightly different from the
177 definition in R2. When there is no evaporative BrC, $E_{\text{abs,950nm}}$ should equal to $E_{\text{abs,370nm}}$, since
178 evaporation of BrC would result in larger $E_{\text{abs,370nm}}$ than $E_{\text{abs,950nm}}$ (Cappa et al., 2012). It is
179 worthy to note that this estimation (R3) actually represents the net result of evaporation and
180 production of BrC in measured aerosol fraction.

182 **3.2 An improved AAE-based method developed by this study**

183 As far as the wavelength segregated AAE is concerned (Sandradewi et al., 2008; Mogo et
184 al., 2012; Utry et al., 2014), the AAE determined over the wavelength of 370–520 nm and 880–
185 950 nm and their diurnal variations were compared in Fig. 2. $\text{AAE}_{370-520}$ were 1.36 ± 0.25 and
186 1.41 ± 0.22 , and $\text{AAE}_{880-950}$ were 0.70 ± 0.07 and 0.68 ± 0.08 for ambient and heated particles,
187 respectively. $\text{AAE}_{880-950}$ were quite similar, indicating there was a thin coating and negligible
188 lensing effect, which could be supported by the chemical composition of the ambient and
189 heated BC-containing particles detected by the SPAMS (Fig. S2). It can be seen that $\text{AAE}_{880-950}$
190 varied in a narrow range (0.63–0.78, 10th–90th of the observed data), and significant difference

191 ($p < 0.001$) was found between $AAE_{370-520}$ and $AAE_{880-950}$ for both ambient and heated
192 particles, respectively. It reflects that σ_{abs} becomes more sensitive to spectral variations at lower
193 wavelengths than at higher wavelengths, which is attributed to the presence of BrC (Mogo et al.,
194 2012). While AAE is sensitive to these microphysical properties (Moosmüller et al., 2009;
195 Lack and Cappa, 2010), the influence of particle size on the AAE for heated aerosols should be
196 limited in the present study, since the difference of hourly average $AAE_{880-950}$ for aerosols
197 before and after heated was in a small range ($2\% \pm 2\%$). Furthermore, there is no significant
198 correlation ($p > 0.05$) between $AAE_{880-950}$ and mass fraction ratio of available coating
199 materials ($PM_{2.5}$ -BC) to BC, which further indicates limited influence of coating materials on
200 the $AAE_{880-950}$. In contrast, the difference between $AAE_{370-520}$ for aerosols before and after
201 heated was relatively higher ($-5\% \pm 11\%$), which is more evidently shown in their diurnal
202 variation (Fig. 2b). Correlation coefficients (r^2) between $AAE_{370-520}$ and $AAE_{590-950}$, $AAE_{660-950}$,
203 and $AAE_{880-950}$ are 0.56, 0.48 and 0.17, respectively, implying limited influence of BrC on the
204 $AAE_{880-950}$. Overall, these results supported that $AAE_{880-950}$ may well represent the AAE of BC
205 dominant aerosols in the present study. “Pure” BC AAE (i.e., $AAE_{880-950}$) of 0.70 ± 0.07
206 inferred in this study is significantly lower than the widely adopted value ($AAE = 1$) (Bond and
207 Bergstrom, 2006). While it might not typical to have AAE of the “pure” BC lower than one, the
208 retrieved value in this study is quite similar to that (AAE between the wavelength at 532 and
209 781 nm) estimated for the “pure” BC dominated from traffic emission in the same region (Yuan
210 et al., 2016). Additionally, lower AAE values (< 1) have also been reported for the “pure” BC in
211 traffic and biomass burning emissions (Kirchstetter et al., 2004), and at urban site in Beijing

212 (Wu et al., 2015). Theoretical calculation shows that AAE decreases with the increasing of
213 “pure” BC sizes (Wu et al., 2015). Regarding the reported volume equivalent diameter peak of
214 “pure” BC to be around 200 nm in this region (Huang et al., 2011), the corresponding
215 theoretical AAE would most probably vary between 0.4 and 1 (Wu et al., 2015). In addition, the
216 fractal like structure of combustion generated BC could also be one possible explanation of the
217 smaller value of AAE measured in this study (Moosmüller et al., 2009; Lack and Cappa, 2010).
218 Therefore, based on the assumption that BC is the dominant absorbing species over the near-IR
219 wavelength (Shrestha et al., 2014), it is proposed to apply $AAE = AAE_{880-950}$, instead of $AAE =$
220 1, in R2 to estimate $\sigma_{\text{abs,BrC}}$. Additionally, the input of hourly variation of $AAE_{880-950}$ further
221 accounts for the factors (e.g., sources and atmospheric processing) that have influence on
222 $AAE_{880-950}$ of BC dominant aerosols.

223 The uncertainty associated with this method was estimated through a sensitivity test. The
224 upper and lower bound of hourly average σ_{abs} ($\sigma_{\text{abs}} \pm \text{std}(\sigma_{\text{abs}})$, $\text{std}(\sigma_{\text{abs}})$ refers to the standard
225 deviation of σ_{abs}) were inputted in R1 and R2 to obtain an estimate of the upper and lower
226 bound of $\sigma_{\text{abs,BrC}}$. $\text{Std}(\sigma_{\text{abs}})$ at 1 min time base is assumed to be three times the noise level
227 (0.016 Mm^{-1}) of AE33, and thus hourly based $\text{std}(\sigma_{\text{abs}})$, averaged over 18 mins measurements
228 of ambient aerosols could be calculated as: $3 \times \sqrt{18} \times 0.016 \text{ Mm}^{-1} = 0.2 \text{ Mm}^{-1}$. The upper
229 and lower bounds of $\sigma_{\text{abs,BrC}}$ was calculated to be 28.8 and 27.1 Mm^{-1} on average, respectively,
230 which is within 3% difference compared to the mean value of $\sigma_{\text{abs,BrC}}$ (28.0 Mm^{-1}).

231 Based on the improved AAE based method, the temporal profiles of the estimated $\sigma_{\text{abs,BrC}}$
232 in Guangdong Atmospheric Supersite as well as the concentration of potassium, the

233 correlation between $\sigma_{\text{abs,BrC}}$ and BC concentration, and the frequency histogram of the
234 $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ are shown in Fig. 3. The estimated $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ distributed within a range
235 of 30%–50%, with an average value at $37.4\% \pm 7.1\%$. Significant correlation ($r^2 = 0.72$)
236 between the estimated $\sigma_{\text{abs,BrC}}$ and BC demonstrates their similar sources (Andreae and
237 Gelencser, 2006; Moosmüller et al., 2009; Saleh et al., 2013; Guo et al., 2014). The
238 correlation between the estimated $\sigma_{\text{abs,BrC}}$ and the concentration of potassium, which is
239 commonly used as a marker for biomass burning activities, was also observed to be
240 significant ($r^2 = 0.38$, $p < 0.001$), indicative of potential contribution of biomass burning to
241 the observed BrC (Lack et al., 2013; Saleh et al., 2014; Hsiao et al., 2016; Rathod et al., 2017)
242 in the PRD region. Potassium originated from sea salt and dust (Pio et al., 2007) was estimated
243 to be $< 5\%$ of the observed potassium in our study. Moderate correlation could be partly
244 explained by that emission of BrC from biomass burning is linked to both burning condition
245 and the emitted organic tracers (Lack et al., 2013; Saleh et al., 2014), and also other sources.

246

247 **3.3 Comparisons of the method developed in this study with the previous** 248 **methods**

249 **3.3.1 Comparison with Method I**

250 It is necessary to note that applying $\text{AAE} = 1$ to anchor absorption for BC and BrC may
251 introduce potential errors, since the AAE for ambient particles (measured between a short and
252 long visible wavelength) has often been observed to be variable, dependent on their physical
253 and chemical properties (Mogo et al., 2012; Lack and Langridge, 2013). Fig. 1 shows the

254 average AAE (1.36 ± 0.25 and 0.70 ± 0.07) of the ambient particles estimated from the linear
255 fit (R1) of σ_{abs} over the wavelength of 370–520 nm and 880–950 nm, respectively.

256 Based on AAE = 1 assumption, $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ was calculated to be $16.8\% \pm 11\%$.
257 While the estimated $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ based on AAE = AAE₈₈₀₋₉₅₀ (the developed method) and
258 AAE = 1 is highly correlated ($r^2 = 0.78$, $p < 0.001$, Fig. 4), the difference can be as large as 20%
259 on average. There were some negative values obtained when assuming AAE = 1, because the
260 observed AAE₃₇₀₋₉₅₀ was lower than one during some periods. Therefore, an arbitrary
261 attribution of AAE to be one may substantially underestimate the $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$.

262

263 3.3.2 Comparison with Method II

264 Based on the Method II, the average estimated $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ over the study was -2.4%
265 $\pm 11.5\%$, which indicates that production of BrC dominated over evaporation, resulted in a net
266 production of BrC when ambient aerosols were heated. Similar estimation in previous studies
267 also resulted in the negative values for $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ (Guo et al., 2014; McMeeking et al.,
268 2014; Nakayama et al., 2014), and it was attributed to the production of BrC when ambient
269 aerosols were heated (Nakayama et al., 2014). A comparison of the $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ estimated
270 by the Methods II and the developed method in Fig. 5 indicates that the production of BrC
271 during heating periods can be substantially important, which may account for $> 40\%$
272 difference between the estimated $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$, and thus should be emphasized when
273 estimating $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ with the Method II. The uncertain evaporation state of BrC during
274 heating might be partly explained by the residue chemical information (Fig. S2). The

275 increased difference during afternoon hours (~50%) (Fig. 5) may further indicate the
276 influence of photochemical activity on the production of BrC during the heating periods. This
277 could be evidenced by negative correlations ($p < 0.001$) between the estimated
278 $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ by Method II and the concentrations of secondary organic carbon (SOC, see
279 SI) and O_3 (Fig. 6). This connection is probably explained by the conversion of SOC to BrC
280 when aerosols were heated, and the conversion should be improved when there are high level
281 of SOC and O_3 . Such fraction of BrC may be attributed to low volatility oxygenated organic
282 compounds, which has also been detected in heated (300 °C) particles (Poulain et al., 2014).
283 Therefore, the estimated $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ based on Method II should suffer substantial error
284 when there is high abundance of SOC.

285 From the co-located SPAMS, it was observed that detected number for internally mixed
286 OC and sulfate (OC-sulfate) particles, which is frequently observed in the atmosphere of PRD
287 region (Zhang et al., 2015), substantially increased during noon hours when aerosols were
288 heated (Fig. S3). More efficient detection of these particles in this case is probably linked to
289 increased absorption of ablation laser by heated OC-sulfate particles (Hatch et al., 2014).
290 With strong absorbing EC cores, the increased absorption is therefore more likely attributed to
291 the formation of absorptive species (i.e., BrC) in heated OC-sulfate particles. Together with
292 significant correlation between the estimated $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ and SOC and sulfate (Fig. 6), it
293 is suspected that the formed BrC during heating measurements might be linked to
294 organosulfate, which was ionized to fragments representative of OC and sulfate in the
295 internally mixed OC and sulfate particles (Hatch et al., 2011). Light absorption at shorter

296 wavelength had been observed for organosulfate (Song et al., 2013). Although secondary
297 organic aerosol (SOA) produced by photo-chemical ageing may contribute to BrC (Lambe et
298 al., 2013; Saleh et al., 2013), our results indicate primary emission plays a dominant role. As
299 discussed above, significant correlations between the estimated $\sigma_{\text{abs,BrC}}$ and BC, and potassium
300 were observed. Furthermore, there is no correlation ($r^2 < 0.01$) between the estimated $\sigma_{\text{abs,BrC}}$
301 by our developed method and SOC (Fig. S4). As noted by Lambe et al. (2013), the extent of
302 absorption by SOA is linked with the magnitude dependent on both precursor type and
303 oxidation level.

304

305 **4 Conclusions**

306 The paper provides an improved AAE-based method to quantitative estimate the light
307 absorption by BrC in real-time based on multi-wavelength light absorption measurements.
308 The analysis shows that the wavelength dependence of BC light absorption can be better
309 approximated over the higher (880–950 nm) wavelengths. Based on simultaneous
310 measurements of σ_{abs} for both ambient and heated aerosols, $\text{AAE}_{880-950}$ was found to show
311 negligible response to the presence of BrC and thus the application of $\text{AAE} = \text{AAE}_{880-950}$ (0.70
312 ± 0.07) for “pure” BC, instead of $\text{AAE} = 1$, should be more acceptable in the estimation of
313 light absorption by BrC in the PRD region. However, the detected AAE value around 0.7 for
314 this work could not represent all kinds of situations, in particular, when BrC absorption over
315 longer wavelength is non-negligible. BrC was estimated to contribute to 37.4% of light
316 absorption at 370 nm by the developed method, and it was largely attributed to primary

317 emissions. The contribution is > 20% on average larger than that estimated by assuming AAE
318 to be one. The results also suggest that the estimated contribution of BrC to total light
319 absorption at 370 nm based on light absorption enhancement method might suffer potential
320 error when there are high loading of SOA. It might serve as evidence for the production of
321 BrC when aerosols were heated, which was neglected in previous studies.

322

323 **Acknowledgements**

324 This work was supported by the National Nature Science Foundation of China (No.
325 41775124), and the State Key Laboratory of Organic Geochemistry (SKLOGA201603A), and
326 Guangdong Provincial Key Laboratory of Environmental Protection and Resources
327 Utilization (2014B030301060). The authors would also like to thank all participants for the
328 sampling at the Guangdong Atmospheric Supersite.

329 **References**

330 Ajtai, T., Filep, A., Utry, N., Schnaiter, M., Linke, C., Bozoki, Z., Szabo, G., Leisner, T.
331 (2011). Inter-comparison of optical absorption coefficients of atmospheric aerosols determined
332 by a multi-wavelength photoacoustic spectrometer and an Aethalometer under sub-urban
333 wintry conditions. *J. Aerosol Sci.* 42: 859-866.

334 Alexander, D. T. L., Crozier, P. A., Anderson, J. R. (2008). Brown carbon spheres in East
335 Asian outflow and their optical properties. *Science* 321: 833-836.

336 Andreae, M. O. and Gelencser, A. (2006). Black carbon or brown carbon? The nature of
337 light-absorbing carbonaceous aerosols. *Atmos. Chem. Phys.* 6: 3131-3148.

338 Arnott, W. P., Hamasha, K., Moosmuller, H., Sheridan, P. J., Ogren, J. A. (2005). Towards
339 aerosol light-absorption measurements with a 7-wavelength Aethalometer: Evaluation with a
340 photoacoustic instrument and 3-wavelength nephelometer. *Aerosol Sci. Tech.* 39: 17-29.

341 Bahadur, R., Praveen, P. S., Xu, Y. Y., Ramanathan, V. (2012). Solar absorption by
342 elemental and brown carbon determined from spectral observations. *Proc. Natl. Acad. Sci. USA*
343 109: 17366-17371.

344 Bergstrom, R. W., Pilewskie, P., Russell, P. B., Redemann, J., Bond, T. C., Quinn, P. K.,
345 Sierau, B. (2007). Spectral absorption properties of atmospheric aerosols. *Atmos. Chem. Phys.*
346 7: 5937-5943.

347 Bi, X. H., Dai, S. H., Zhang, G. H., Qiu, N., Li, M., Wang, X. M., Chen, D. H., Peng, P. A.,
348 Sheng, G. Y., Fu, J. M., Zhou, Z. (2015). Real-time and single-particle volatility of elemental
349 carbon-containing particles in the urban area of Pearl River Delta region, China. *Atmos.*
350 *Environ.* 118: 194-202.

351 Bond, T. C. and Bergstrom, R. W. (2006). Light absorption by carbonaceous particles: An
352 investigative review. *Aerosol Sci. Tech.* 40: 27-67.

353 Cappa, C. D., Onasch, T. B., Massoli, P., Worsnop, D. R., Bates, T. S., Cross, E. S.,
354 Davidovits, P., Hakala, J., Hayden, K. L., Jobson, B. T., Kolesar, K. R., Lack, D. A., Lerner, B.
355 M., Li, S. M., Mellon, D., Nuaaman, I., Olfert, J. S., Petaja, T., Quinn, P. K., Song, C.,

356 Subramanian, R., Williams, E. J., Zaveri, R. A. (2012). Radiative Absorption Enhancements
357 Due to the Mixing State of Atmospheric Black Carbon. *Science* 337: 1078-1081.

358 Chen, X. J., Lai, S. C., Gao, Y., Zhang, Y. Y., Zhao, Y., Chen, D. H., Zheng, J. Y., Zhong,
359 L. J., Lee, S. C., Chen, B. (2016). Reconstructed Light Extinction Coefficients of Fine
360 Particulate Matter in Rural Guangzhou, Southern China. *Aerosol Air Qual. Res.* 16: 1981-1990.

361 Cheng, Y., He, K. B., Zheng, M., Duan, F. K., Du, Z. Y., Ma, Y. L., Tan, J. H., Yang, F. M.,
362 Liu, J. M., Zhang, X. L., Weber, R. J., Bergin, M. H., Russell, A. G. (2011). Mass absorption
363 efficiency of elemental carbon and water-soluble organic carbon in Beijing, China. *Atmos.*
364 *Chem. Phys.* 11: 11497-11510.

365 Cheng, Y. F., Wiedensohler, A., Eichler, H., Su, H., Gnauk, T., Brüggemann, E.,
366 Herrmann, H., Heintzenberg, J., Slanina, J., Tuch, T., Hu, M., Zhang, Y. H. (2008). Aerosol
367 optical properties and related chemical apportionment at Xinken in Pearl River Delta of China.
368 *Atmos. Environ.* 42: 6351-6372.

369 Corr, C. A., Hall, S. R., Ullmann, K., Anderson, B. E., Beyersdorf, A. J., Thornhill, K. L.,
370 Cubison, M. J., Jimenez, J. L., Wisthaler, A., Dibb, J. E. (2012). Spectral absorption of biomass
371 burning aerosol determined from retrieved single scattering albedo during ARCTAS. *Atmos.*
372 *Chem. Phys.* 12: 10505-10518.

373 Drinovec, L., Močnik, G., Zotter, P., Prévôt, A. S. H., Ruckstuhl, C., Coz, E., Rupakheti,
374 M., Sciare, J., Müller, T., Wiedensohler, A., Hansen, A. D. A. (2015). The "dual-spot"
375 Aethalometer: an improved measurement of aerosol black carbon with real-time loading
376 compensation. *Atmos. Meas. Tech.* 8: 1965-1979.

377 Favez, O., Alfaro, S. C., Sciare, J., Cachier, H., Abdelwahab, M. M. (2009). Ambient
378 measurements of light-absorption by agricultural waste burning organic aerosols. *J. Aerosol Sci.*
379 40: 613-620.

380 Feng, Y., Ramanathan, V., Kotamarthi, V. R. (2013). Brown carbon: a significant
381 atmospheric absorber of solar radiation? *Atmos. Chem. Phys.* 13: 8607-8621.

382 Guo, X. S., Nakayama, T., Yamada, H., Inomata, S., Tonokura, K., Matsumi, Y. (2014).
383 Measurement of the light absorbing properties of diesel exhaust particles using a
384 three-wavelength photoacoustic spectrometer. *Atmos. Environ.* 94: 428-437.

385 Gyawali, M., Arnott, W. P., Lewis, K., Moosmuller, H. (2009). In situ aerosol optics in
386 Reno, NV, USA during and after the summer 2008 California wildfires and the influence of
387 absorbing and non-absorbing organic coatings on spectral light absorption. *Atmos. Chem. Phys.*
388 9: 8007-8015.

389 Harrison, R. M., Beddows, D. C. S., Hu, L., Yin, J. (2012). Comparison of methods for
390 evaluation of wood smoke and estimation of UK ambient concentrations. *Atmos. Chem. Phys.*
391 12: 8271-8283.

392 Hatch, L. E., Creamean, J. M., Ault, A. P., Surratt, J. D., Chan, M. N., Seinfeld, J. H.,
393 Edgerton, E. S., Su, Y., Prather, K. A. (2011). Measurements of Isoprene-Derived
394 Organosulfates in Ambient Aerosols by Aerosol Time-of-Flight Mass Spectrometry - Part 1:
395 Single Particle Atmospheric Observations in Atlanta. *Environ. Sci. Technol.* 45: 5105-5111.

396 Hatch, L. E., Pratt, K. A., Huffman, J. A., Jimenez, J. L., Prather, K. A. (2014). Impacts of
397 Aerosol Aging on Laser Desorption/Ionization in Single-Particle Mass Spectrometers. *Aerosol*
398 *Sci. Tech.*: 1050-1058.

399 Hsiao, T. C., Ye, W. C., Wang, S. H., Tsay, S. C., Chen, W. N., Lin, N. H., Lee, C. T.,
400 Hung, H. M., Chuang, M. T., Chantara, S. (2016). Investigation of the CCN Activity, BC and
401 UVBC Mass Concentrations of Biomass Burning Aerosols during the 2013 BASELInE
402 Campaign. *Aerosol Air Qual. Res.* 16: 2742-2756.

403 Huang, X. F., Gao, R. S., Schwarz, J. P., He, L. Y., Fahey, D. W., Watts, L. A.,
404 McComiskey, A., Cooper, O. R., Sun, T. L., Zeng, L. W., Hu, M., Zhang, Y. H. (2011). Black
405 carbon measurements in the Pearl River Delta region of China. *J. Geophys. Res.* 116: 445-451.

406 Kirchstetter, T. W., Novakov, T., Hobbs, P. V. (2004). Evidence that the spectral
407 dependence of light absorption by aerosols is affected by organic carbon. *J. Geophys.*
408 *Res.-Atmos.* 109: 21208.

409 Kirchstetter, T. W. and Thatcher, T. L. (2012). Contribution of organic carbon to wood
410 smoke particulate matter absorption of solar radiation. *Atmos. Chem. Phys.* 12: 6067-6072.

411 Lack, D. A. and Cappa, C. D. (2010). Impact of brown and clear carbon on light absorption
412 enhancement, single scatter albedo and absorption wavelength dependence of black carbon.
413 *Atmos. Chem. Phys.* 10: 4207-4220.

414 Lack, D. A., Bahreni, R., Langridge, J. M., Gilman, J. B., Middlebrook, A. M. (2013).
415 Brown carbon absorption linked to organic mass tracers in biomass burning particles. *Atmos.*
416 *Chem. Phys.* 13: 2415-2422.

417 Lack, D. A. and Langridge, J. M. (2013). On the attribution of black and brown carbon
418 light absorption using the Angstrom exponent. *Atmos. Chem. Phys.* 13: 10535-10543.

419 Lambe, A. T., Cappa, C. D., Massoli, P., Onasch, T. B., Forestieri, S. D., Martin, A. T.,
420 Cummings, M. J., Croasdale, D. R., Brune, W. H., Worsnop, D. R., Davidovits, P. (2013).
421 Relationship between Oxidation Level and Optical Properties of Secondary Organic Aerosol.
422 *Environ. Sci. Technol.* 47: 6349-6357.

423 Laskin, A., Laskin, J., Nizkorodov, S. A. (2015). Chemistry of Atmospheric Brown
424 Carbon. *Chem. Rev.* 115: 4335-4382.

425 Li, L., Huang, Z. X., Dong, J. G., Li, M., Gao, W., Nian, H. Q., Fu, Z., Zhang, G. H., Bi, X.
426 H., Cheng, P., Zhou, Z. (2011). Real time bipolar time-of-flight mass spectrometer for
427 analyzing single aerosol particles. *Intl. J. Mass. Spectrom.* 303: 118-124.

428 Lin, C. C., Yang, L. S., Cheng, Y. H. (2016). Ambient PM_{2.5}, Black Carbon, and Particle
429 Size-Resolved Number Concentrations and the Angstrom Exponent Value of Aerosols during
430 the Firework Display at the Lantern Festival in Southern Taiwan. *Aerosol Air Qual. Res.* 16:
431 373-387.

432 Liu, J., Scheuer, E., Dibb, J., Ziemba, L. D., Thornhill, K. L., Anderson, B. E., Wisthaler,
433 A., Mikoviny, T., Devi, J. J., Bergin, M., Weber, R. J. (2014). Brown carbon in the continental
434 troposphere. *Geophys. Res. Lett.* 41: 2191-2195.

435 Liu, J., Scheuer, E., Dibb, J., Diskin, G. S., Ziemba, L. D., Thornhill, K. L., Anderson, B.
436 E., Wisthaler, A., Mikoviny, T., Devi, J. J., Bergin, M., Perring, A. E., Markovic, M. Z.,

437 Schwarz, J. P., Campuzano-Jost, P., Day, D. A., Jimenez, J. L., Weber, R. J. (2015). Brown
438 carbon aerosol in the North American continental troposphere: sources, abundance, and
439 radiative forcing. *Atmos. Chem. Phys.* 15: 7841-7858.

440 McMeeking, G. R., Fortner, E., Onasch, T. B., Taylor, J. W., Flynn, M., Coe, H.,
441 Kreidenweis, S. M. (2014). Impacts of nonrefractory material on light absorption by aerosols
442 emitted from biomass burning. *J. Geophys. Res.-Atmos.* 119: 12272-12286.

443 Mogo, S., Cachorro, V. E., de Frutos, A., Rodrigues, A. (2012). Absorption Angstrom
444 exponents of aerosols and light absorbing carbon (LAC) obtained from in situ data in Covilha,
445 central Portugal. *J. Environ. Monitor.* 14: 3174-3181.

446 Moosmüller, H., Chakrabarty, R. K., Arnott, W. P. (2009). Aerosol light absorption and its
447 measurement: A review. *J. Quant. Spectrosc. Ra-diat. Transfer* 110: 844-878.

448 Moosmüller, H., Chakrabarty, R. K., Ehlers, K. M., Arnott, W. P. (2011). Absorption
449 Ångström coefficient, brown carbon, and aerosols: basic concepts, bulk matter, and spherical
450 particles. *Atmos. Chem. Phys.* 11: 1217-1225.

451 Nakayama, T., Ikeda, Y., Sawada, Y., Setoguchi, Y., Ogawa, S., Kawana, K., Mochida,
452 M., Ikemori, F., Matsumoto, K., Matsumi, Y. (2014). Properties of light-absorbing aerosols in
453 the Nagoya urban area, Japan, in August 2011 and January 2012: Contributions of brown
454 carbon and lensing effect. *J. Geophys. Res.-Atmos.* 119: 12721-12739.

455 Pintér, M., Ajtai, T., Kiss-Albert, G., Kiss, D., Utry, N., Janovszky, P., Palásti, D., Smausz,
456 T., Kohut, A., Hopp, B., Galbács, G., Kukovecz, Á., Kónya, Z., Szabó, G., Bozóki, Z. (2018).
457 Thermo-optical properties of residential coals and combustion aerosols. *Atmos. Environ.* 178:
458 118-128.

459 Poulain, L., Birmili, W., Canonaco, F., Crippa, M., Wu, Z. J., Nordmann, S., Spindler, G.,
460 Prevot, A. S. H., Wiedensohler, A., Herrmann, H. (2014). Chemical mass balance of 300
461 degrees C non-volatile particles at the tropospheric research site Melpitz, Germany. *Atmos.*
462 *Chem. Phys.* 14: 10145-10162.

463 Qin, Y. M., Tan, H. B., Li, Y. J., Schurman, M. I., Li, F., Canonaco, F., Prevot, A. S. H.,
464 Chan, C. K. (2017). Impacts of traffic emissions on atmospheric particulate nitrate and organics

465 at a downwind site on the periphery of Guangzhou, China. *Atmos. Chem. Phys.* 17:
466 10245-10258.

467 Rathod, T., Sahu, S. K., Tiwari, M., Yousaf, A., Bhangare, R. C., Pandit, G. G. (2017).
468 Light Absorbing Properties of Brown Carbon Generated from Pyrolytic Combustion of
469 Household Biofuels. *Aerosol Air Qual. Res.* 17: 108-116.

470 Rumsey, I. C., Cowen, K. A., Walker, J. T., Kelly, T. J., Hanft, E. A., Mishoe, K., Rogers,
471 C., Proost, R., Beachley, G. M., Lear, G., Frelink, T., Otjes, R. P. (2014). An assessment of the
472 performance of the Monitor for AeRosols and GAses in ambient air (MARGA): a
473 semi-continuous method for soluble compounds. *Atmos. Chem. Phys.* 14: 5639-5658.

474 Saleh, R., Hennigan, C. J., McMeeking, G. R., Chuang, W. K., Robinson, E. S., Coe, H.,
475 Donahue, N. M., Robinson, A. L. (2013). Absorptivity of brown carbon in fresh and
476 photo-chemically aged biomass-burning emissions. *Atmos. Chem. Phys.* 13: 7683-7693.

477 Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R.
478 C., Presto, A. A., Dubey, M. K., Yokelson, R. J., Donahue, N. M., Robinson, A. L. (2014).
479 Brownness of organics in aerosols from biomass burning linked to their black carbon content.
480 *Nature Geosci.* 7: 647-650.

481 Sandradewi, J., Prevot, A. S. H., Weingartner, E., Schmidhauser, R., Gysel, M.,
482 Baltensperger, U. (2008). A study of wood burning and traffic aerosols in an Alpine valley
483 using a multi-wavelength Aethalometer. *Atmos. Environ.* 42: 101-112.

484 Shamjad, P. M., Tripathi, S. N., Pathak, R., Hallquist, M., Arola, A., Bergin, M. H. (2015).
485 Contribution of Brown Carbon to Direct Radiative Forcing over the Indo-Gangetic Plain.
486 *Environ. Sci. Technol.* 49: 10474-10481.

487 Shrestha, R., Kim, S. W., Yoon, S. C., Kim, J. H. (2014). Attribution of aerosol light
488 absorption to black carbon and volatile aerosols. *Environ. Monit. Assess.* 186: 4743-4751.

489 Song, C., Gyawali, M., Zaveri, R. A., Shilling, J. E., Arnott, W. P. (2013). Light
490 absorption by secondary organic aerosol from alpha-pinene: Effects of oxidants, seed aerosol
491 acidity, and relative humidity. *J. Geophys. Res.-Atmos.* 118: 11741-11749.

492 Stevanovic, S., Miljevic, B., Madl, P., Clifford, S., Ristovski, Z. (2015). Characterisation
493 of a Commercially Available Thermodenuder and Diffusion Drier for Ultrafine Particles
494 Losses. *Aerosol Air Qual. Res.* 15: 357-+.

495 Utry, N., Ajtai, T., Filep, Á., Pintér, M., Török, Z., Bozóki, Z., Szabó, G. (2014).
496 Correlations between absorption Angström exponent (AAE) of wintertime ambient urban
497 aerosol and its physical and chemical properties. *Atmos. Environ.* 91: 52-59.

498 Villani, P., Picard, D., Marchand, N., Laj, P. (2007). Design and validation of a 6-volatility
499 tandem differential mobility analyzer (VTDMA). *Aerosol Sci. Tech.* 41: 898-906.

500 Wu, Y. F., Yan, P., Tian, P., Tao, J., Li, L., Chen, J. M., Zhang, Y. M., Cao, N. W., Chen,
501 C., Zhang, R. J. (2015). Spectral Light Absorption of Ambient Aerosols in Urban Beijing
502 during Summer: An Intercomparison of Measurements from a Range of Instruments. *Aerosol*
503 *Air Qual. Res.* 15: 1178-1187.

504 Yang, M., Howell, S. G., Zhuang, J., Huebert, B. J. (2009). Attribution of aerosol light
505 absorption to black carbon, brown carbon, and dust in China - interpretations of atmospheric
506 measurements during EAST-AIRE. *Atmos. Chem. Phys.* 9: 2035-2050.

507 Yuan, J. F., Huang, X. F., Cao, L. M., Cui, J., Zhu, Q., Huang, C. N., Lan, Z. J., He, L. Y.
508 (2016). Light absorption of brown carbon aerosol in the PRD region of China. *Atmos. Chem.*
509 *Phys.* 16: 1433-1443.

510 Zhang, G. H., Bi, X. H., Chan, L. Y., Wang, X. M., Sheng, G. Y., Fu, J. M. (2013a).
511 Size-segregated chemical characteristics of aerosol during haze in an urban area of the Pearl
512 River Delta region, China. *Urban Climate* 4: 74-84.

513 Zhang, G. H., Bi, X. H., He, J. J., Chen, D. H., Chan, L. Y., Xie, G. W., Wang, X. M.,
514 Sheng, G. Y., Fu, J. M., Zhou, Z. (2014). Variation of secondary coatings associated with
515 elemental carbon by single particle analysis. *Atmos. Environ.* 92: 162-170.

516 Zhang, G. H., Han, B. X., Bi, X. H., Dai, S. H., Huang, W., Chen, D. H., Wang, X. M.,
517 Sheng, G. Y., Fu, J. M., Zhou, Z. (2015). Characteristics of individual particles in the
518 atmosphere of Guangzhou by single particle mass spectrometry. *Atmos. Res.* 153: 286-295.

519 Zhang, X., Lin, Y.-H., Surratt, J. D., Weber, R. J. (2013b). Sources, Composition and
520 Absorption Ångström Exponent of Light-absorbing Organic Components in Aerosol Extracts
521 from the Los Angeles Basin. *Environ. Sci. Technol.* 47: 3685-3693.

522 Zhu, J., Crozier, P. A., Anderson, J. R. (2013). Characterization of light-absorbing carbon
523 particles at three altitudes in East Asian outflow by transmission electron microscopy. *Atmos.*
524 *Chem. Phys.* 13: 6359-6371.

525

ACCEPTED MANUSCRIPT

526 List of figure captions

527 Fig. 1. An illustration of the method to estimate the contribution of light
528 absorption from BrC, based on the assumption of AAE = 1 for Method I, and AAE =
529 $AAE_{880-950}$ for our developed method.

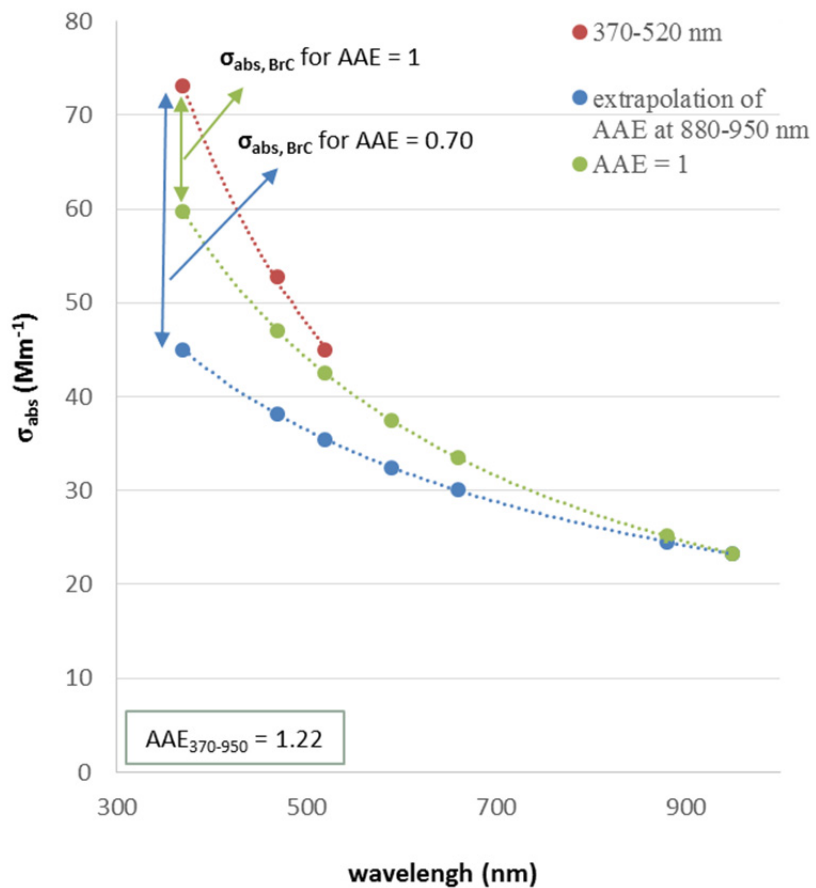
530 Fig. 2. (a) Comparison of AAE at the wavelength of 370–520, 880–950, and
531 370–950 nm, and diurnal variations of AAE at the wavelength of (b) 370–520 nm and
532 (c) 880–950 nm, for aerosols before and after heated, respectively.

533 Fig. 3. (a) Temporal profile of $\sigma_{\text{abs,BrC}}$ and the concentration of potassium, (b)
534 correlation analysis of $\sigma_{\text{abs,BrC}}$ and the concentration of BC, and (c) the histogram
535 summarizing of the $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ over this study.

536 Fig. 4. Comparison of $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ estimated by our developed method and
537 Methods I based on AAE = 1.

538 Fig. 5. Comparisons of (a) diurnal variations and (b) hourly averaged
539 $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ estimated by Methods II and our developed method.

540 Fig. 6. Correlation plots between the $\sigma_{\text{abs,BrC}}/\sigma_{\text{abs,370nm}}$ estimated by Method II
541 and the concentrations of O₃, SOC, OC and sulfate.



SCRIPT

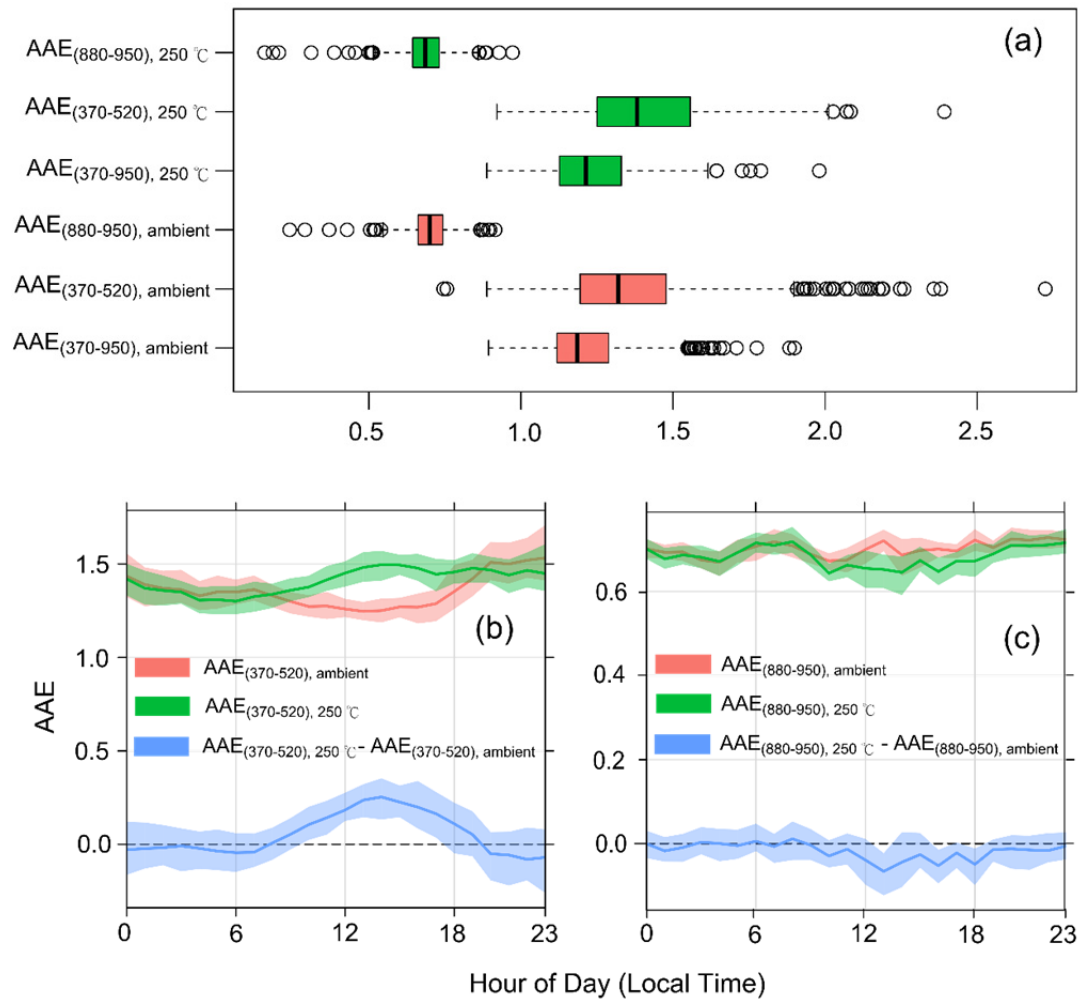
542

543

544

Fig. 1.

ACCEPTED MANUSCRIPT



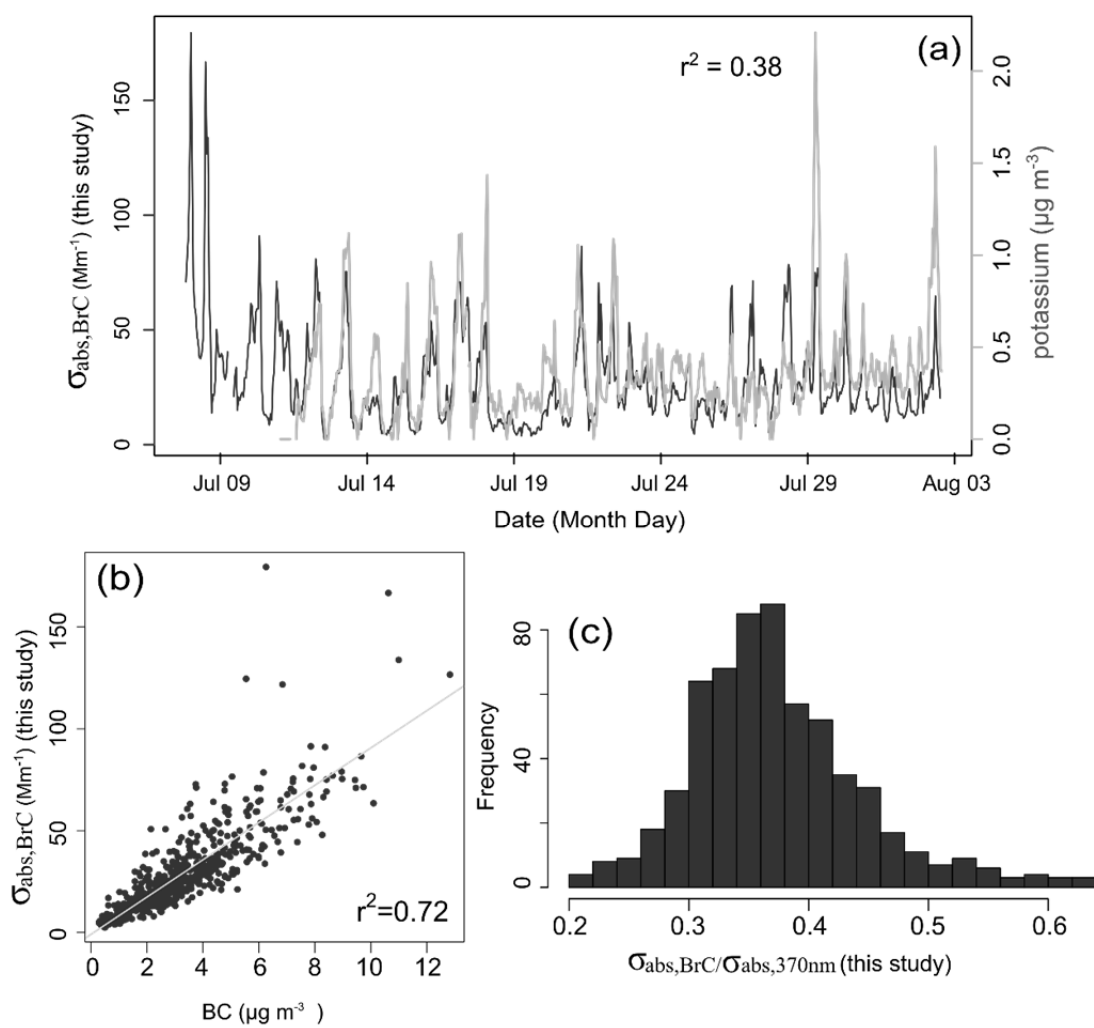
545

546

547

Fig. 2.

ACCEPTED



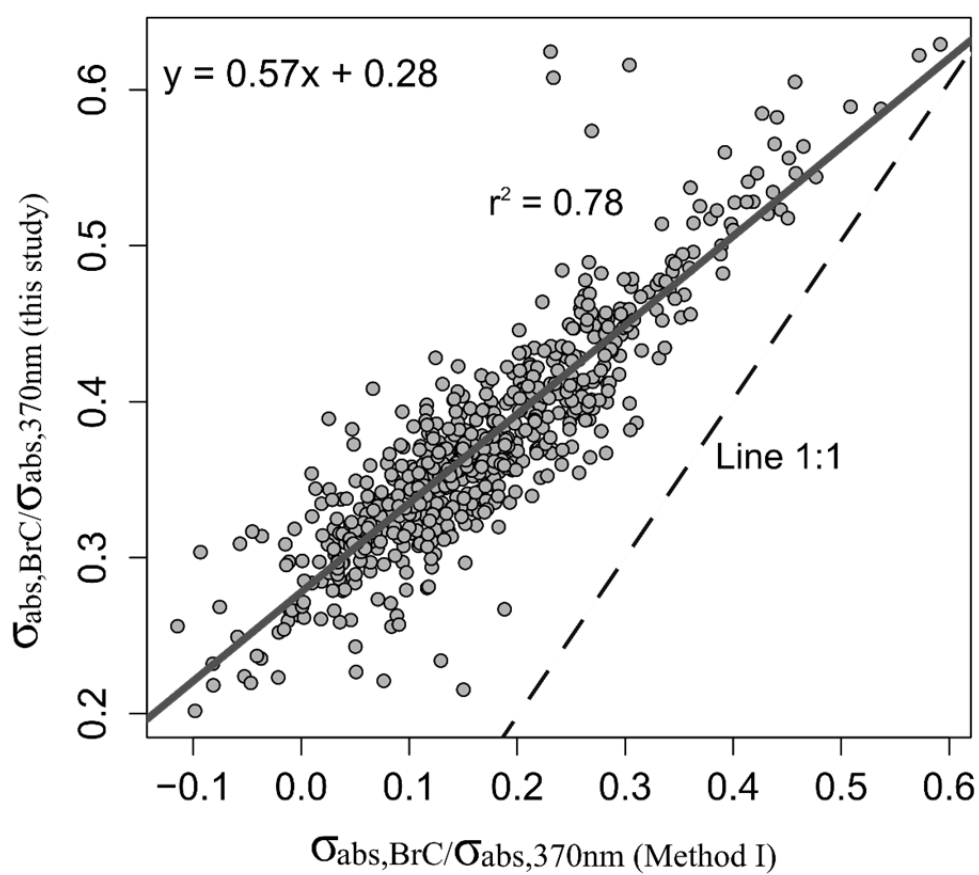
548

549

550

Fig. 3.

ACCEPTED



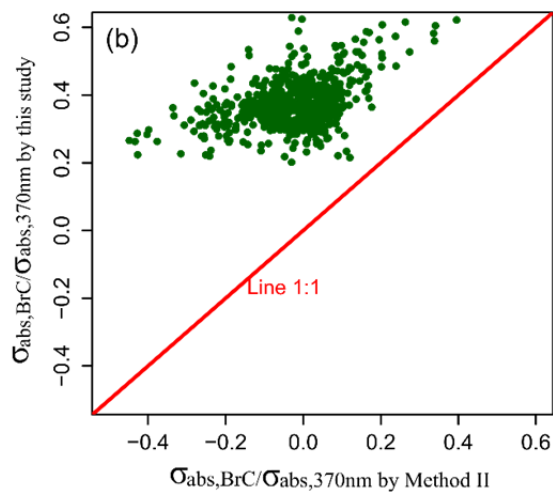
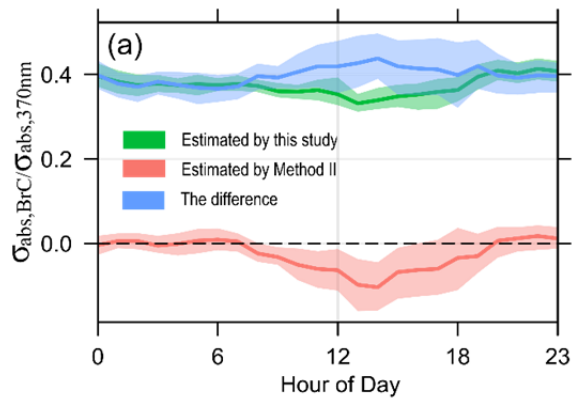
551

552

553

Fig. 4.

ACCEPTED

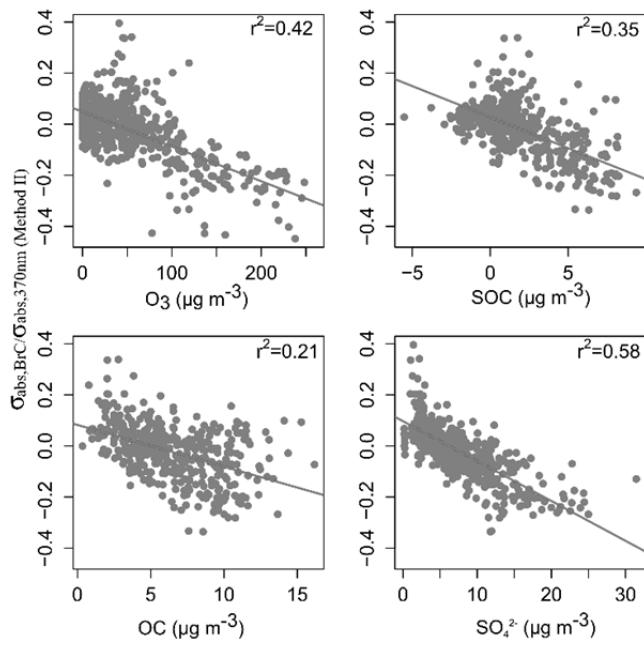


554

555

556 Fig. 5.

557



558

559

560 Fig. 6.

ACCEPTED MANUSCRIPT